

# Linking soil N<sub>2</sub>O emissions with soil microbial community abundance and structure related to nitrogen cycle in two acid forest soils

Hongling Qin · Xiaoyi Xing · Yafang Tang ·  
Haijun Hou · Jie Yang · Rong Shen · Wenzhao Zhang ·  
Yi Liu · Wenxue Wei

Received: 3 May 2018 / Accepted: 23 October 2018  
© Springer Nature Switzerland AG 2018

## Abstract

**Aims** Tree species and seasonal change influence N<sub>2</sub>O flux and microbial communities, but the mechanisms are unclear. We studied N<sub>2</sub>O flux in soils planted with slash

---

Hongling Qin and Xiaoyi Xing contributed equally to this research.

## Highlights

- Tree species sharply affected *in situ* N<sub>2</sub>O flux and functional microbe community
- Abundance and structure of nitrifiers and denitrifiers varied between tree species depending on soil substrate availability
- Community structure of nitrifiers and denitrifiers was more affected by tree species than by seasonal change
- Microbial community composition rather than abundance determined N<sub>2</sub>O flux in acid forest soils

---

Responsible Editor: Elizabeth M Baggs.

**Electronic supplementary material** The online version of this article (<https://doi.org/10.1007/s11104-018-3863-7>) contains supplementary material, which is available to authorized users.

H. Qin · X. Xing · H. Hou · J. Yang · R. Shen · W. Zhang ·  
Y. Liu · W. Wei (✉)

Key Laboratory of Agro-ecological Processes in Subtropical Regions, Taoyuan Agro-ecosystem Research Station, Institute of Subtropical Agriculture, Chinese Academy of Sciences, Changsha 410125, China  
e-mail: wenxuewei@isa.ac.cn

X. Xing · Y. Tang  
University of Chinese Academy of Sciences, Beijing 100000, China

Y. Tang  
Hubei Key Laboratory of Quality Control of Characteristic Fruits and Vegetables, College of Life Science and Technology, Hubei Engineering University, Xiaogan 432000, China

pine and oil-seed camellia trees. We sampled on typical days of the four seasons. We tested whether N-cycling communities respond more to tree species or seasonal change. We assessed how tree species affect N<sub>2</sub>O flux.

**Methods** We used qPCR and RFLP to determine abundance and community composition of *amoA*-containing bacteria (AOB) and archaea (AOA), and denitrifiers that contain the *narG*, *nirK*, *nirS*, and *nosZ* genes.

**Results** N<sub>2</sub>O flux rate and soil characteristics varied significantly between forest soils and sampling seasons. Abundance of all detected genes, but not of the *nirS* gene, was significantly affected by tree species. Differences in gene abundance between days in different seasons were found only for *narG*, *nirK*, and *nosZ*. Functional microbial community composition in the soil varied between the tree species for most of the genes studied, but varied, not significantly, slightly among sampling days. Differences in the abundance and community composition of nitrifiers and denitrifiers between tree species depended on soil concentration of NH<sub>4</sub><sup>+</sup>, NO<sub>3</sub><sup>-</sup>, and dissolved organic carbon (DOC). N<sub>2</sub>O flux rate was affected by community composition, but not abundance of nitrifiers and denitrifiers. Temperature, NO<sub>3</sub><sup>-</sup>, and DOC concentrations significantly affected N<sub>2</sub>O flux.

**Conclusions** Tree species influenced N<sub>2</sub>O flux more than seasonal change, by altering community composition and environmental factors rather than nitrifier/denitrifier abundance.

**Keywords** N<sub>2</sub>O emission · Nitrification · Denitrification · Acid forest soil · Tree species · Seasonal change

## Introduction

Nitrous oxide (N<sub>2</sub>O) is a powerful and enduring greenhouse gas whose warming effect in the troposphere is 265-times stronger than that of carbon dioxide, on a 100-year time horizon (IPCC 2014). Further, N<sub>2</sub>O reacts with ozone in the stratosphere and is a dominant compound that threatens to deplete ozone in the twenty-first century (Ravishankara et al. 2009). N<sub>2</sub>O emissions from temperate forest soils amount to 1.0 (0.2–2.0) Tg N y<sup>-1</sup> (Morishita et al. 2011). Thus, strategies to mitigate N<sub>2</sub>O emissions should take forest soils into account.

There have been some artificial afforestation experiments to study how tree species affect N<sub>2</sub>O flux, but the results are contradictory. A study on artificial afforestation in Siberia predicted that deciduous species produce more N<sub>2</sub>O than conifers, because of a large discrepancy between N<sub>2</sub>O production and reduction rates under deciduous species (Menyailo et al. 2002). In contrast, Oura et al. (2001) reported that coniferous stands emit more N<sub>2</sub>O than adjacent deciduous stands that were affected by nitrogen deposition. A study of a wide range of European forest soils revealed that N<sub>2</sub>O emissions were highest in deciduous forests, but these forests were also situated on more fertile soils (Ambus et al. 2006). Further, Liu and Greaver (2009) found that soil N status rather than forest type could be the main factor determining the magnitude of N<sub>2</sub>O emissions. The role of tree species in modulating N<sub>2</sub>O flux, therefore, remains unclear.

In general, microbiological processes of nitrification and denitrification regulate the production and consumption of N<sub>2</sub>O, while both processes are influenced by soil temperature, moisture content, pH, and carbon and nitrogen content (Brown et al. 2012). Microbial ammonia oxidation is the first and also the rate-limiting step in nitrification. The functional gene *amoA*, which codes for ammonia monooxygenase, is used as a biomarker for ammonia-oxidising bacteria (AOB) and ammonia-oxidising archaea (AOA). Several genes have been used as molecular markers of the denitrification steps: *narG* (which codes for membrane-bound nitrate reductase); *nirK* (for copper-containing nitrite reductase); *nirS* (for Cyt cd1-containing nitrite reductase); and *nosZ* (for nitrous oxide reductase) (Zhang et al. 2016). The contribution of nitrification and denitrification to N<sub>2</sub>O emissions depends on climate and soil conditions (Wolf and Brumme 2002). Seasonal changes in soil temperature and moisture may cause shifts in

N<sub>2</sub>O emissions and in the microbial communities involved in nitrification and denitrification. The N<sub>2</sub>O flux increases with increasing soil moisture when denitrifying microorganisms are promoted (Menyailo and Hungate 2015). N<sub>2</sub>O emissions increase with increasing soil temperature, because the rates of enzymatic processes increase with temperature, if other factors are not limiting (Schindlbacher et al. 2004). In forest ecosystems, soil temperature and moisture do not change simultaneously. The influence of seasonal change on N<sub>2</sub>O emissions and on related functional microbial communities is therefore a complex phenomenon.

In the subtropical red soil hilly areas in China, the planting area of oil-seed camellia has increased annually because of its economic importance (Luo et al. 2016). Slash pine is an important tree species for afforestation because of drought resistance and high survival rate (Xu et al. 2018). Previously, we documented that tree species affects soil biological fertility (Zheng et al. 2006): natural secondary forest and tea-oil plantations had higher soil biological fertility than Chinese fir and slash pine plantations. In that study, we ascribed this difference to the higher litter production and quality, higher root biomass, richer plant species, better soil ecological condition and faster plant growth of the secondary forests and tea-oil plantations. Soil fertility drives the effect of vegetative cover on N<sub>2</sub>O emissions (Guardia et al. 2016). Thus, we expect different tree species to affect soil N<sub>2</sub>O production differently.

In this work, we assessed in situ N<sub>2</sub>O flux, soil environmental factors, abundance and microbial (genetic) community composition involved in nitrification and denitrification in two soils of oil-seed camellia forests and slash pine forests, from 2012 to 2014, and over the four seasons. Our objectives were 1) to investigate the seasonal change in N<sub>2</sub>O flux in soils planted with different tree species; 2) to test whether N-cycling communities varied with tree species and seasonal change; and 3) to explain how tree species affect N<sub>2</sub>O flux. We hypothesized that tree species would induce a shift in in situ N<sub>2</sub>O flux; that soils from oil-seed camellia forests would have higher N<sub>2</sub>O flux rates than soils from slash pine forests; and that tree species would induce specific soil characteristics that affect the abundance and community composition of nitrification and denitrification microbes, thereby influencing the in situ N<sub>2</sub>O flux.

## Materials and methods

### Experimental site

The sampling site is located at the Taoyuan Agroecosystem Research Station (111°26' E, 28°55' N and altitude 92.2–125.3 m) and represents a typical hilly area of the Hunan Province, China. The soil is developed from quaternary red clay, and the region is characterized by a subtropical humid monsoon climate, with annual average air temperature of 16.5 °C, annual average precipitation of 1448 mm, daily average sunshine duration of 15 h 13 min, and an annual frost-free period of 283 days (Qin et al. 2013).

Soil samples were collected from slash pine (*Pinus elliottii*) and oil-seed camellia (*Camellia oleifera* Abel.) forests in a long-term land use experiment at the Taoyuan Agroecosystem Research Station. The experiment, with eight land use plots of 62 m × 20 m each, was established in 1995. The field has a 15° slope, a length of 62 m, and an area of 10,000 m<sup>2</sup>. Each land use plot extended from the top to bottom of the sloped field. The field was too small to allow replicated plots. At the start of the experiment, the slash pine and oil-seed camellia forest plots had the same soil properties, were free of human disturbance, and the trees were of the same age. Soil properties at the start of the experiment in 1995 were pH (H<sub>2</sub>O), 4.45; soil organic matter (SOM), 28.9 g kg<sup>-1</sup>; and total N, 1.93 g kg<sup>-1</sup>. The similarity of initial site conditions provided an adequate background for comparison of the tree-mediated effects on soils (Binkley and Giardina 1998). Thus, we assumed that any soil differences between two corresponding treatments revealed by the study at the same site would be an effect of tree species. Soil properties in the two forests are summarized in Table 1. As can be seen, in 2012, higher SOM and total N was registered in soils of the oil-seed camellia forest.

Soil samples were collected on typical days of the four seasons: winter, spring, summer, and autumn. We chose typical days based on temperature (Supplementary Fig. 1). In 2012 we sampled on 20 October. In 2013, we sampled on 17 January, 12 April, 9 July, and 22 October. In 2014 we sampled on 19 January, 3 April, and 19 July. The days in October were typical days of autumn, those in January of winter, those in April of spring, and those in July of summer, respectively. We defined Oct 2012–July 2013 as year 1 and Oct 2013–July 2014 as year 2. For each plot in the slash pine and

oil-seed camellia forests, three replicate sampling sites (50 m × 5 m) were selected at the same altitude across the land use plot. At each sampling site, ten soil cores along a “S” shaped line were collected (0–20 cm) and mixed thoroughly. In all, 48 soil samples were collected (2 species × 4 seasons × 2 years × 3 replicates). From each sample, a sub-sample (approximately 100 g) was packed into a sterile bag, which was immediately stored in liquid nitrogen in the field. Each sub-sample was then freeze dried (Neocool, Yamato Scientific Co., Ltd., Tokyo, Japan) in the lab under sterile conditions, ground to a fine powder after removal of root residues, and stored at –80 °C for subsequent molecular analysis. The remaining soil from each sample (approximately 400 g) was stored at 4 °C for chemical analysis.

For all samples, we analysed soil moisture, soil temperature and soil substrate availability of nitrate-N (NO<sub>3</sub><sup>-</sup>-N), ammonia-N (NH<sub>4</sub><sup>+</sup>-N), and dissolved organic carbon (DOC). Soil moisture was determined by the oven drying method. Soil temperature at 5 cm soil depth was measured using a geothermometer. Soil NO<sub>3</sub><sup>-</sup>-N and NH<sub>4</sub><sup>+</sup>-N were extracted with 1 M KCl and analysed with a continuous-flow injection analyser (FIAstar 5000, Foss Corporation, Hillerod, Denmark). DOC was extracted using 0.5 M K<sub>2</sub>SO<sub>4</sub> and analysed for total oxidizable carbon (TOC-VMP, Shimadzu Corporation, Kyoto, Japan).

### N<sub>2</sub>O emission measurement

In situ N<sub>2</sub>O gas fluxes were monitored by the static opaque chamber gas-chromatography (GC) method according to Zheng et al. (2008). The chamber (50 cm long × 50 cm wide × 20 cm high) was composed of insulated stainless-steel sheet metal and was equipped with a fan to ensure thorough gas mixing. The base was inserted 10 cm into the soil 2 months before the commencement of gas sampling to minimize the effects of disturbance. The static opaque chamber was equipped in the middle of each soil sampling site. On the same day of soil sampling, gases were sampled in the same time from each land use plot of slash pine and oil-seed camellia forests. Each sampling site, a 30-mL samples were collected using gas-tight syringes. We sampled 5 times per site on each day, at 15 min intervals from 9:00 to 10:00, when the air temperature is as close as possible to the mean daily temperature (Zhang et al. 2012). Concentrations were determined using a gas chromatograph (Agilent 7890A, Agilent Technologies, Santa

**Table 1** Basic properties of the two forest soils used in this study of soil N<sub>2</sub>O emissions

Tree species	pH (H <sub>2</sub> O)	SOM (g kg <sup>-1</sup> )	Alkaline N (mg kg <sup>-1</sup> )	Olsen P (mg kg <sup>-1</sup> )	available K (mg kg <sup>-1</sup> )	Total-N (g kg <sup>-1</sup> )	Total P (g kg <sup>-1</sup> )	Total K (g kg <sup>-1</sup> )
Oil-seed camellia	4.33	32.93	139.45	3.30	46.82	1.61	0.31	9.83
Slash pine	4.31	27.22	125.83	2.53	62.13	1.33	0.32	12.68

Clara, CA, USA) fitted with an electron capture detector (ECD) for N<sub>2</sub>O analyses at 350 °C. The in situ N<sub>2</sub>O fluxes were determined by linear regression and expressed in µg N m<sup>-2</sup> h<sup>-1</sup>.

#### DNA extraction

Soil microbial DNA was extracted from 0.5 g soil stored at -80 °C, as described by Chen et al. (2010). DNA quality and concentration were measured using a NanoDrop NA-1000 spectrophotometer (Nanodrop Technologies, Wilmington, DE, USA). The extractions were replicated three times, pooled, and stored at -80 °C for further analysis.

#### Quantitative PCR

Primer pairs AOB-*amoA*-1F/AOB-*amoA*-2R, AOA-*amoA*-23F/AOA-*amoA*-616R, *narG*-571F/*narG*-773R, *nirK*-876F/*nirK*-1055R, *nirS*-cd3aF/*nirS*-R3cd, and *nosZ*-1126F/*nosZ*-1381R were used for quantifying gene abundance (Supplementary Table 1). The qPCR assay was performed with an ABI 7900HT (Applied Biosystems, Foster City, CA, USA). The 10 µL reaction mixture consisted of 5 µL 2× SYBR green mix II (TaKaRa Biotechnology Co. Ltd., Dalian, China), 0.2 µL 50× Rox Reference Dye (TaKaRa Biotechnology Co. Ltd., Dalian, China), 0.4 µL each 10 µM forward and reverse primers, and 5 ng DNA template. A standard curve was obtained from serial 10× dilutions of linearized plasmids containing the target gene. Three parallel PCR replicates of all samples were performed on each plate. A melting curve analysis was conducted following the assay to verify the specificity of the amplification product. The PCR efficiency was in the range of 90–110%.

#### T-RFLP

The terminal-restriction fragment length polymorphism (T-RFLP) method was used for community composition

analysis. Amplification of AOB *amoA* gene sequences from soil microbial DNA was done using a nested polymerase chain reaction (PCR) protocol. First-round PCR primers used were *amoA*-2F/5R and second-round PCR primers *amoA*-1F/2R. The primer pairs *amoA*-23F/616R, *narG*-145F/773R, *nirK*-517F/1055R, *nirS*-cd3aF/R3cd and *nosZ*-2002F/2002R were used for PCR amplification of AOA *amoA*, *narG*, *nirK*, *nirS* and *nosZ* genes (Supplementary Table 1). Forward primers were labelled with 6-carboxyfluorescein at the 5' ends, except for *amoA*-2F. Approximately 200 ng of each amplicon was digested with 5 U restriction enzymes; TaqI were used for AOA *amoA*, AOB *amoA*, *narG*, *nirK* fragments, HaeIII and MspI for *nirS* and *nosZ* fragments, respectively. There are 11 T-RFs for AOA *amoA*, 28 T-RFs for AOB *amoA*, 26 T-RFs for *narG*, 29 T-RFs for *nirK*, 37 T-RFs for *nirS* and 13 T-RFs for *nosZ* genes.

#### Statistical analyses

Two-way analysis of variance (ANOVA) with repeated measures of 2 years was used to determine significant differences in the interaction of tree species and sampling seasons. The copy numbers of all functional genes were log-transformed, and the normality of all data was checked before running the ANOVA test. The data were found to be normally distributed. One-way ANOVA with repeated measures was used to evaluate the effect of tree species and seasons on the soil properties and abundance of genes, since two-way ANOVA did not reveal any significant interaction between tree species and sampling seasons for most parameters measured. Significant differences between treatments were identified by the least significant difference (LSD) test at the 5% level. Pearson correlation and regression analysis were conducted to determine the relationships among N<sub>2</sub>O flux, abundance of N-associated functional genes, and soil parameters using all data from the 48 plots. All statistical analyses were performed using SPSS

10. Principal component analysis (PCA) based on the T-RFLP data was used to assess the similarities between the nitrifying and denitrifying microbial community structures in the slash pine and oil-tea camellia soils) among days from four seasons. In addition, redundancy analysis (RDA) was used to evaluate the effects of soil properties on microbial community composition. A Monte Carlo test with 999 random permutations was used to test for significant associations between environmental factors (soil temperature, moisture,  $\text{NO}_3^-$ -N,  $\text{NH}_4^+$ -N, DOC and  $\text{N}_2\text{O}$  flux) and species (nitrifying and denitrifying microbial community composition). PCA and RDA were performed using Canoco v.4.5 for Windows.

## Results

### In situ $\text{N}_2\text{O}$ flux rate and soil properties

Significant differences were detected for in situ  $\text{N}_2\text{O}$  flux rate and for all measured soil characteristics, except soil temperature, between the oil-seed camellia and slash pine soils. In addition, sampling days from different seasons also had significant effects on all soil properties and in situ  $\text{N}_2\text{O}$  flux rate (Table 2). In situ  $\text{N}_2\text{O}$  fluxes ranged from 3.78 to  $596.16 \mu\text{g N m}^{-2} \text{h}^{-1}$ , with average values of 128.89 and  $30.87 \mu\text{g N m}^{-2} \text{h}^{-1}$  in oil-seed camellia and slash pine soils, respectively. In situ  $\text{N}_2\text{O}$  flux rate, soil moisture, and soil  $\text{NH}_4^+$ -N and DOC content were higher in the soil of the oil-seed camellia forest than that of the slash pine forest. However,  $\text{NO}_3^-$ -N content revealed an opposite trend (Table 3). In situ  $\text{N}_2\text{O}$  flux rate varied among days from different seasons, with the highest rate observed in summer days and lowest in autumn days. Soil temperature ranged from 6.62 to  $24.64 \text{ }^\circ\text{C}$  and varied strongly among sampling days from different seasons. Soil moisture, soil  $\text{NO}_3^-$ -N, soil  $\text{NH}_4^+$ -N and soil DOC were higher in summer days and in autumn days, but lower in spring (Table 4).

### Abundance of nitrification and denitrification functional genes

There was significant difference in functional gene abundance between the oil-seed camellia and slash pine soils, except for the *nirK* gene. However,

**Table 2** Two-way analysis of variance of  $\text{N}_2\text{O}$  flux, soil environmental factors, and gene abundances

Variable	df	$\text{N}_2\text{O}$ flux	Temperature	Moisture	$\text{NO}_3^-$ -N	$\text{NH}_4^+$ -N	DOC	AOA <i>amoA</i>	AOB <i>amoA</i>	<i>narG</i>	<i>nirK</i>	<i>nirS</i>	<i>nosZ</i>
Tree species	F ratio	1975.49	0.42	9.08	220.09	108.89	51.57	20.88	9.11	4.74	0.21	15.86	15.23
	<i>p</i>	<b>0.00</b>	0.52	<b>0.01</b>	<b>0.00</b>	<b>0.00</b>	<b>0.00</b>	<b>0.00</b>	<b>0.00</b>	<b>0.04</b>	0.65	<b>0.00</b>	<b>0.00</b>
Season	F ratio	3	1742.32	3.27	32.67	4.90	10.67	2.30	0.90	6.47	4.49	2.18	3.25
	<i>p</i>	<b>0.00</b>	<b>0.00</b>	<b>0.03</b>	<b>0.00</b>	<b>0.01</b>	<b>0.00</b>	0.10	0.45	<b>0.00</b>	<b>0.01</b>	0.11	<b>0.03</b>
Tree species $\times$ season	F ratio	3	706.24	1.50	26.80	8.52	1.85	1.71	0.13	0.19	0.78	1.20	1.24
	<i>p</i>	<b>0.00</b>	0.06	0.23	<b>0.00</b>	<b>0.00</b>	0.16	0.18	0.94	0.90	0.51	0.32	0.31

AOA *amoA*, AOB *amoA*, *narG*, *nirK*, *nirS*, and *nosZ* are the copy numbers of these genes. And the copy numbers of genes were  $\log_{10}$  transformed; Bold typing indicates significance at 5%

**Table 3** One-way table of means of soil properties and abundance of genes between two tree species. Abundances of genes were log<sub>10</sub> transformed

Items		Slash pine	Oil-seed camelli
The abundance of genes (copies per g dried soil)	<i>AOA amoA</i>	7.33 (0.42) a	6.75 (0.76) b
	<i>AOB amoA</i>	6.27 (0.41) b	6.83 (0.83) a
	<i>narG</i>	8.93 (0.44) a	8.67 (0.55) b
	<i>nirK</i>	7.31 (0.42) a	7.36 (0.54) a
	<i>nirS</i>	6.35 (0.59) a	5.78 (0.51) b
	<i>nosZ</i>	7.55 (0.35) b	7.87 (0.50) a
Soil moisture (% by weight)		21.53 (4.47) b	23.62 (5.17) a
Soil NH <sub>4</sub> <sup>+</sup> -N (mg kg <sup>-1</sup> )		7.32 (2.66) b	14.86 (6.58) a
Soil NO <sub>3</sub> <sup>-</sup> -N (mg kg <sup>-1</sup> )		6.40 (3.69) a	1.66 (0.63) b
Soil DOC (mg kg <sup>-1</sup> )		305.23 (50.68) b	375.81 (74.69) a
Soil temperature (°C)		15.88 (6.47) a	15.77 (6.93) a
N <sub>2</sub> O flux rate (μg N m <sup>-2</sup> h <sup>-1</sup> )		30.87 (30.52) b	128.89 (183.06) a

Numbers within parenthesis are standard deviations,  $n = 24$ ; Different letters within rows indicate significant difference between tree species ( $p < 0.05$ )

significant differences between sampling days from different seasons were detected only for denitrification functional genes of *narG*, *nirK*, and *nosZ* (Table 2). The abundance of *AOA amoA*, *narG*, and *nirS* genes was higher in soils of the slash pine forest than in those of the oil-seed camellia forest, while the opposite was true for *AOB amoA* and *nosZ* genes. (Table 3). The abundance of the denitrifying genes was higher in the days of winter than in any other season, with average abundance values of  $1.10 \times 10^9$ ,  $5.42 \times 10^7$  and  $7.39 \times 10^7$  copies g<sup>-1</sup> dry soil, for *narG*, *nirK* and *nosZ*, respectively, in the soil of the oil-seed camellia forest. In contrast, no difference was found in soils of the slash pine forest (Fig. 1). We also detected a positive relationship between the abundance of genes *narG* and *nirK* and soil moisture content (Table 5).

#### Community composition of nitrification and denitrification functional genes

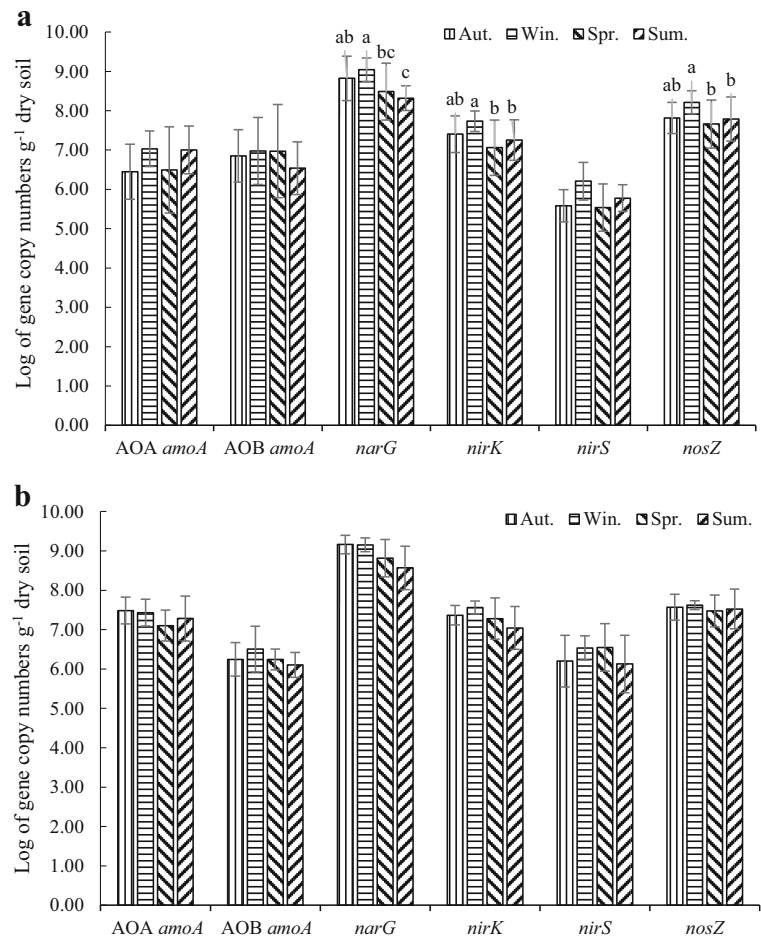
The average relative abundance of T-RFs for *AOA amoA*, *AOB amoA*, *narG*, *nirK*, *nirS*, and *nosZ* genes is shown in Fig. 2. We detected a difference in community composition between the soils of the two forest types. However, soil microbial community composition of all detected genes showed no significant difference among days from different seasons. In most instances, T-RFs were greater in the oil-seed forest than in the slash pine forest soils: these included dominant T-RFs of 54 bp for *AOA amoA* and 282 bp for *AOB amoA* (by 30.56% and 25.94%, respectively); four dominant T-RFs, of 206, 222, 246, and 260 bp for *narG* (by 12.58%, 18.64%, 8.11%, and 7.62%, respectively);

**Table 4** One-way table of means of soil properties for four seasons

Season	N <sub>2</sub> O flux (μg N m <sup>-2</sup> h <sup>-1</sup> )	Soil temperature (°C)	Soil moisture (% by weight)	Soil NO <sub>3</sub> <sup>-</sup> -N (mg kg <sup>-1</sup> )	Soil NH <sub>4</sub> <sup>+</sup> -N (mg kg <sup>-1</sup> )	Soil DOC (mg kg <sup>-1</sup> )
Autumn	22.40 (28.17) d	17.85 (0.84) b	23.66 (8.27) a	5.61 (4.77) a	11.38 (7.14) a	302.02 (39.52) bc
Winter	53.44 (43.08) b	6.62 (1.08) d	22.02 (2.63) bc	2.59 (1.61) b	12.75 (8.62) a	353.18 (53.17) a
Spring	32.48 (10.36) c	14.21 (0.32) c	21.08 (2.13) c	2.32 (0.93) b	8.90 (3.78) b	329.82 (108.67) b
Summer	211.21 (232.60) a	24.64 (1.22) a	23.55 (4.20) ab	5.61 (4.09) a	11.31 (4.40) a	377.06 (53.85) a

Numbers within parenthesis are standard deviations,  $n = 12$ ; Different letters within columns indicate significant difference among four seasons ( $p < 0.05$ )

**Fig. 1** Abundance of AOA *amoA*, AOB *amoA*, *narG*, *nirK*, *nirS*, and *nosZ* genes in soils under different forest vegetation of oil-seed camellia (a) and slash pine (b). Spr., spring; Sum., summer; Aut., Autumn; Win., winter, the same below. Error bars are standard deviations for seasons. Different letters above bars indicate significant differences among seasons ( $P < 0.05$ )



two dominant T-RFs of 190 and 224 bp for *nirK* (by 12.37% and 10.33%, respectively); three dominant T-RFs of 116, 136, and 210 bp for *nirS* (by 8.14%, 11.60%, and 16.32%, respectively); and one dominant T-RF of 350 bp for *nosZ*, by 9.75%. The dominant T-RFs of 130 bp for AOA *amoA* and 280 bp for AOB *amoA* were less, by 33.95% and 27.35%, respectively, in the oil-seed camellia than the slash pine forest soils. One dominant T-RFs of 248 bp for *narG* was less in the oil-seed camellia forest, by 13.71%; this T-RF, and some minor T-RFs for *narG* in soils of the slash pine forest (228 bp, 244 bp, 250 bp, and 266 bp), disappeared from soils of the oil-seed camellia forest. Two dominant T-RFs for *nirK* of 192 bp and 436 bp were less, by 11.89% and 5.07%, respectively, in the oil-seed camellia than in the slash pine forest soils. Four dominant T-RFs for *nirS* of 78 bp, 188 bp, 190 bp, and 406 bp were less, by 7.99%,

7.53%, 18.28%, and 9.60%, respectively, in the oil-seed camellia than in the slash pine forest soils. Two T-RFs for *nosZ* of 206 bp and 220 bp were less, by 8.24% and 9.34%, respectively, in the oil-seed camellia than in the slash pine forest soils.

PCA showed that variations in the nitrification and denitrification microbial community structure were affected more by tree species than by seasonal change (Fig. 3). The first two principal components (PCs) accounted for 98.8 and 1.0%, 98.8 and 1.0%, 76.2 and 16.1%, 47.5 and 28.9%, 70.4 and 17.0%, and 51.7 and 34.8%, respectively, of the total variability recorded in AOA *amoA*, AOB *amoA*, *narG*, *nirK*, *nirS*, and *nosZ* genes. For all the detected genes, except for *nosZ* gene, the main separation of slash pine and oil-seed camellia forests along PC1 reveals that differences between the two tree species caused most of the variation in the data, while the smaller separation of the seasons along PC2 was responsible for less of the variation. The Monte Carlo test based on

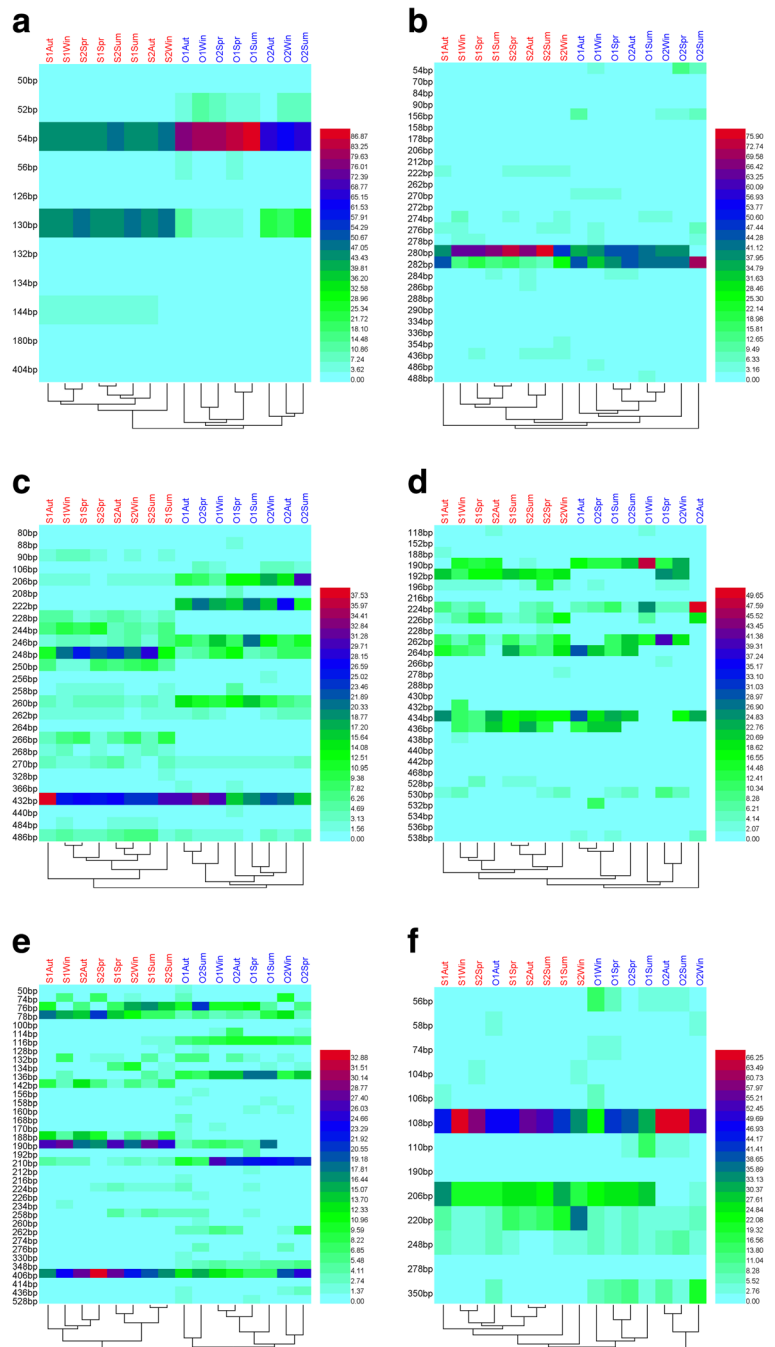
**Table 5** Pearson correlation coefficients for gene abundance and soil properties in two forest soils

Items	AOB <i>amoA</i>	<i>narG</i>	<i>nirK</i>	<i>nirS</i>	<i>nosZ</i>	Soil moisture	Soil NH <sub>4</sub> <sup>+</sup> -N	Soil NO <sub>3</sub> <sup>-</sup> -N	Soil DOC	Soil temperature	N <sub>2</sub> O flux
Oil-seed camelli	0.23	<b>0.47</b>	<b>0.63</b>	<b>0.48</b>	<b>0.54</b>	0.28	0.37	0.21	-0.25	-0.04	0.39
AOB <i>amoA</i>		0.33	<b>0.45</b>	0.34	<b>0.49</b>	0.35	-0.06	-0.22	-0.04	-0.17	-0.07
<i>narG</i>			0.79	<b>0.56</b>	<b>0.71</b>	<b>0.41</b>	0.17	-0.30	<b>-0.40</b>	<b>-0.40</b>	-0.26
<i>nirK</i>				<b>0.72</b>	<b>0.86</b>	<b>0.45</b>	0.30	-0.16	-0.34	-0.27	0.03
<i>nirS</i>					<b>0.64</b>	0.23	0.16	-0.17	-0.25	-0.30	0.07
<i>nosZ</i>					<b>0.41</b>		0.29	-0.05	-0.18	-0.30	0.22
Soil moisture							-0.06	-0.39	-0.11	0.11	0.14
Soil NH <sub>4</sub> <sup>+</sup> -N								0.28	-0.07	-0.19	0.21
Soil NO <sub>3</sub> <sup>-</sup> -N									0.02	0.27	<b>0.54</b>
Soil DOC										0.08	0.33
Soil temperature											<b>0.42</b>
Slash pine											-0.21
AOA <i>amoA</i>	<b>0.49</b>	<b>0.49</b>	<b>0.55</b>	0.20	<b>0.80</b>	<b>0.55</b>	0.20	0.30	-0.27	-0.07	-0.29
AOB <i>amoA</i>		<b>0.44</b>	0.47	0.11	<b>0.48</b>	0.25	0.18	-0.11	-0.03	-0.34	-0.29
<i>narG</i>			<b>0.88</b>	0.21	<b>0.46</b>	0.06	-0.07	-0.10	-0.15	<b>-0.40</b>	<b>-0.66</b>
<i>nirK</i>				0.25	<b>0.62</b>	0.06	0.07	-0.12	-0.21	<b>-0.42</b>	<b>-0.53</b>
<i>nirS</i>					0.04	-0.27	0.19	-0.36	-0.39	-0.29	-0.32
<i>nosZ</i>						<b>0.50</b>	0.10	0.24	-0.24	-0.12	-0.12
Soil moisture							0.05	<b>0.51</b>	-0.06	0.21	0.18
Soil NH <sub>4</sub> <sup>+</sup> -N								-0.14	-0.38	0.19	0.11
Soil NO <sub>3</sub> <sup>-</sup> -N									0.10	<b>0.64</b>	0.28
Soil DOC										0.10	<b>0.41</b>
Soil temperature											<b>0.65</b>

Bold typing indicates significant difference ( $P < 0.05$ ),  $n = 24$



**Fig. 2** Heat map of T-RFLP profiles showing the average relative abundances of AOA *amoA* (a), AOB *amoA* (b), *narG* (c), *nirK* (d) T-RFs with endonuclease TaqI, of *nirS* (e) and *nosZ* (f) T-RFs with endonuclease HaeIII and MspI, respectively. Relative abundances of T-RFs are given as percentages within the graph indicate the size (bp) of the experimental T-RFs as determined by T-RFLP. S, slash pine soil; O, oil-seed camellia soil. 1 and 2 mean the first and second sampling years

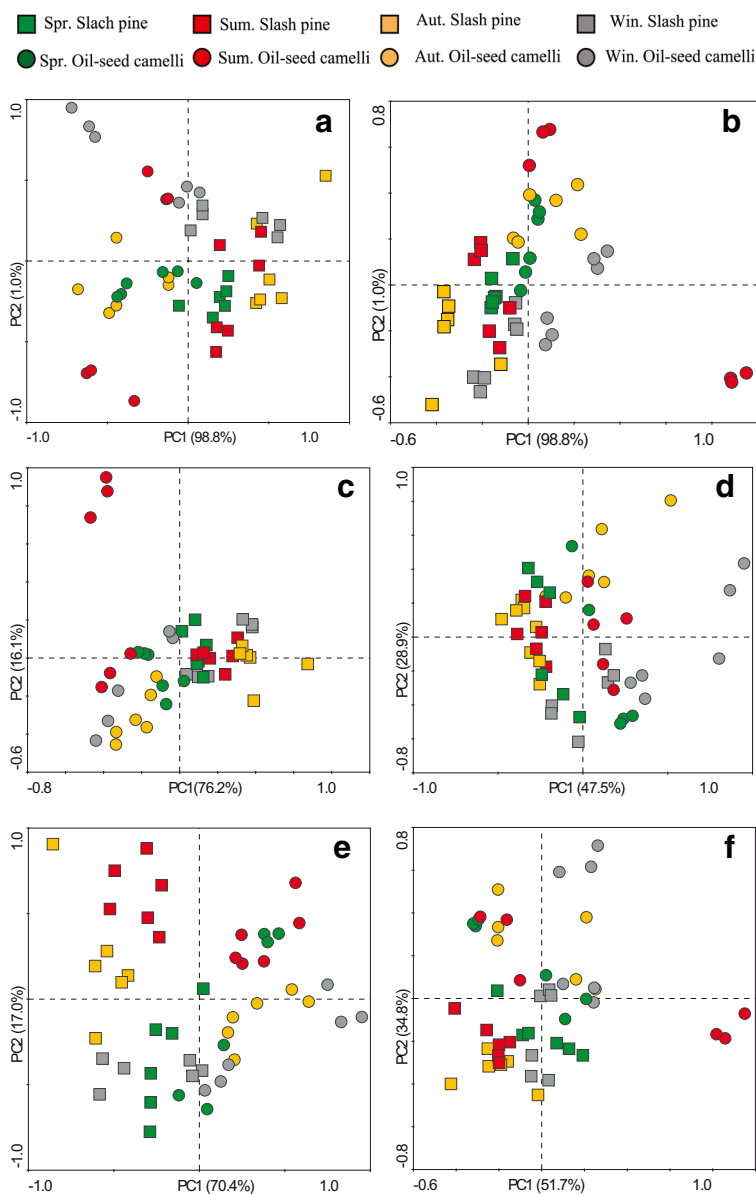


redundancy analysis (RDA) indicated that soil  $\text{NO}_3^-$ -N,  $\text{NH}_4^+$ -N, and DOC significantly affected the nitrifying and denitrifying community compositions, except that DOC did not affect *nosZ*-containing denitrifier composition. Soil temperature and moisture significantly affected *nirK*-containing denitrifiers; and soil moisture also affected AOB *amoA*-containing nitrifiers (Table 6).

Relationships among  $\text{N}_2\text{O}$  flux, soil environmental factors, functional gene abundance and community composition

Pearson correlation coefficient analysis revealed that in situ  $\text{N}_2\text{O}$  flux rate was significantly and positively correlated with soil temperature and soil  $\text{NO}_3^-$ -N in soils of the

**Fig. 3** Principal coordinate analysis based on the T-RFLP profile of AOA *amoA* (a), AOB *amoA* (b), *narG* (c), *nirK* (d), *nirS* (e) and *nosZ* (f) genes



**Table 6** Monte Carlo test of redundancy analysis (RDA) based on the T-RFLP profile of nitrification and denitrification genes and soil properties

Genes	Soil temperature		Soil moisture		Soil NO <sub>3</sub> <sup>-</sup> -N		Soil NH <sub>4</sub> <sup>+</sup> -N		Soil DOC		N <sub>2</sub> O flux	
	F	<i>p</i>	F	<i>p</i>	F	<i>p</i>	F	<i>p</i>	F	<i>p</i>	F	<i>p</i>
AOA <i>amoA</i>	0.76	0.15	2.32	0.13	30.19	<b>0.00</b>	15.97	<b>0.00</b>	14.82	<b>0.00</b>	1.53	0.26
AOB <i>amoA</i>	0.53	0.52	5.27	<b>0.02</b>	7.87	<b>0.01</b>	8.78	<b>0.00</b>	7.78	<b>0.02</b>	26.09	<b>0.00</b>
<i>narG</i>	0.46	0.74	1.41	0.21	15.29	<b>0.00</b>	11.80	<b>0.00</b>	8.27	<b>0.00</b>	7.98	<b>0.00</b>
<i>nirK</i>	4.93	<b>0.00</b>	2.39	<b>0.03</b>	5.07	<b>0.01</b>	11.18	<b>0.00</b>	7.29	<b>0.00</b>	2.50	<b>0.04</b>
<i>nirS</i>	2.25	0.08	1.16	0.26	12.80	<b>0.00</b>	10.03	<b>0.00</b>	8.89	<b>0.00</b>	2.70	<b>0.03</b>
<i>nosZ</i>	1.49	0.18	0.64	0.62	4.33	<b>0.02</b>	5.72	<b>0.00</b>	2.33	0.07	6.08	<b>0.00</b>

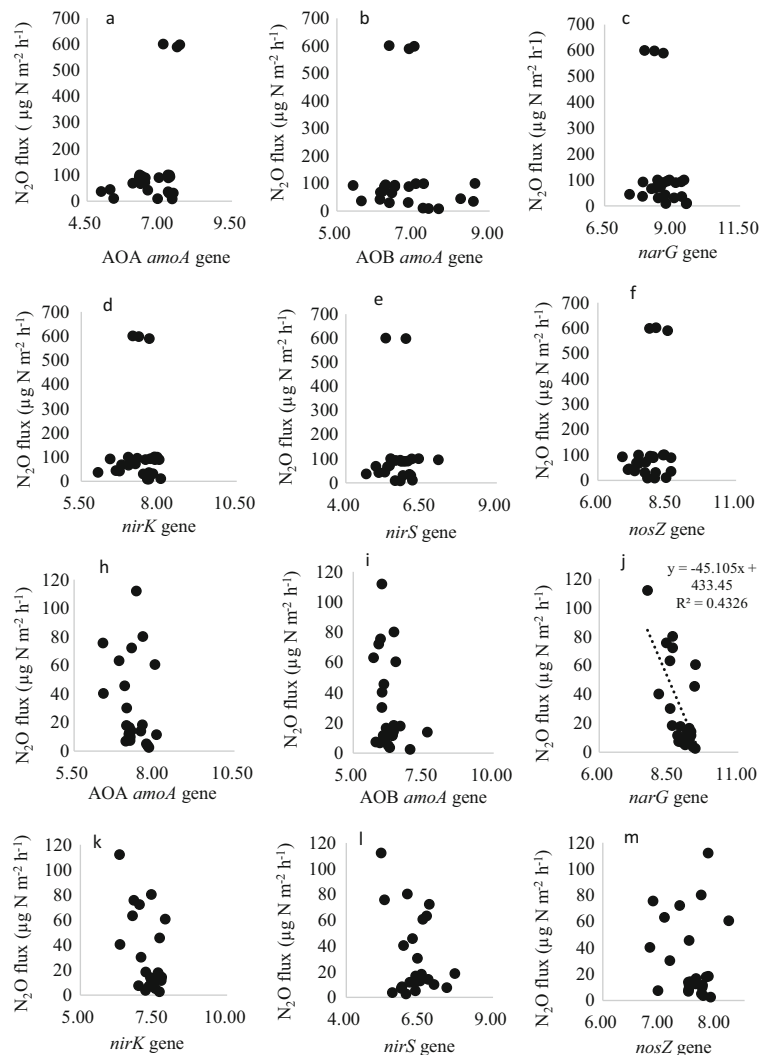
Bold typing indicates significant difference ( $p < 0.05$ )

oil-seed camellia forest, and with soil temperature and DOC in soils of the slash pine forest (Table 5). Nevertheless, no significant positive relationship between in situ  $N_2O$  flux rate and abundance of nitrification and denitrification functional genes could be found in the soils of either forest type based on regression analysis (Fig. 4). However, the Monte Carlo test showed that in situ  $N_2O$  flux rate was significantly affected by all detected community compositions of all detected genes, except for AOA *amoA* gene (Table 6). Furthermore, correspondingly with  $N_2O$  flux in situ, the average relative abundance of some T-RFs of all detected genes in soils of the oil-seed camellia forest significantly exceeded that in soils of the slash pine forest, regardless of seasonal change (Fig. 2).

## Discussion

Tree species influenced  $N_2O$  flux, with higher flux values in oil-seed camellia than slash pine forest soils, supporting our hypotheses. This confirmed that forests with deciduous species should produce more  $N_2O$  than forests with conifers. Slash pine is a conifer with needle-like leaves that contain large amounts of substances that are difficult to decompose, such as tannins, resins, cellulose, and lignin. In contrast, oil-seed camellia is a deciduous species, which produces leaf litter containing substances easily decomposed by bacteria, such as nitrogen and water (Yang et al. 2014). Zheng et al. (2006) found that oil-seed camellia forest soil had higher vegetation biomass and root biomass, and therefore, higher

**Fig. 4** Regression analysis between  $N_2O$  flux and abundance of AOA *amoA* (a), AOB *amoA* (b), *narG* (c), *nirK* (d), *nirS* (e), and *nosZ* (f) genes in oil-seed camellia soil; and AOA *amoA* (h), AOB *amoA* (i), *narG* (g), *nirK* (k), *nirS* (l), and *nosZ* (m) genes in slash pine soil. AOA *amoA* gene, AOB *amoA* gene, *narG* gene, *nirK* gene, *nirS* gene, and *nosZ* gene are the log<sub>10</sub> transformed copy numbers per g dried soil of AOA *amoA*, AOB *amoA*, *narG*, *nirK*, *nirS* and *nosZ* genes



SOM content, than did slash pine forest soil. Likewise, we found that soils of the oil-seed camellia forest had higher soil moisture,  $\text{NH}_4^+\text{-N}$ , and DOC, but lower soil  $\text{NO}_3^-\text{-N}$ , than soils of the slash pine forest. Soil environmental factors (temperature, concentration of  $\text{NO}_3^-$ , and DOC) were significantly related to  $\text{N}_2\text{O}$  fluxes. This indicates that tree species induced shifts in soil environmental factors, which were largely responsible for the variation in  $\text{N}_2\text{O}$  fluxes in acidic forest soils.

Tree species also influenced the microbial community related to  $\text{N}_2\text{O}$  emission. The differences in the abundance and community composition of nitrifiers and denitrifiers between tree species depended on soil substrate availability. The higher abundance of the AOB *amoA* gene in the soil of the oil-seed camellia forest is consistent with the higher content of  $\text{NH}_4^+\text{-N}$  in this soil. However, the AOA *amoA* gene showed the opposite trend to AOB *amoA* gene. AOB and AOA have been found to co-exist in most soils, both mediating the first step of autotrophic nitrification (Schleper 2010). The relative roles of AOB and AOA have been debated since Leininger et al. (2006) first reported the dominance of AOA in 12 pristine agricultural soils from three climatic zones: they reported that AOB population was more sensitive to soil  $\text{NH}_4^+\text{-N}$  than AOA in acidic forest soils. Those findings were supported by the findings of Yang et al. (2016), who observed that AOB were more responsive than AOA to nitrogen source in an agricultural soil. Conversely, in our study, higher abundance of the *narG* and *nirS* genes in the soils of the slash pine forest was consistent with the higher  $\text{NO}_3^-\text{-N}$  content. Although both *nirK* and *nirS* encodes nitrite reductase, *nirK* abundance was not significantly different between tree species. Levy-Booth et al. (2014) reported that organisms that contain the *nirS* gene often do not contain the *nirK* gene. Our study suggested that the effect of tree species on populations of nitrite reducers was associated with the *nirS* gene, but not with the *nirK* gene. Finally, higher abundance of the *nosZ* gene in the soils of the oil-seed camellia forest was consistent with the higher  $\text{N}_2\text{O}$  flux rate in our study, consistent with the  $\text{N}_2\text{O}$  concentration in the surface soil (Wang et al. 2017). We found that the composition of all detected nitrifiers and denitrifiers varied between the two forest types, which were affected by soil  $\text{NO}_3^-\text{-N}$ ,  $\text{NH}_4^+\text{-N}$ , and DOC availability.

Production of  $\text{N}_2\text{O}$  is essentially mediated by microbial functional groups. Interestingly, data from this study revealed that community composition (not the abundance of nitrification and denitrification genes) was related to in situ  $\text{N}_2\text{O}$  flux in acidic forest soils. A few studies have found significant correlations between  $\text{N}_2\text{O}$  emissions and *narG*, *nirS*, and *nirK* gene numbers under favourable conditions for denitrification. For example, rainfall or irrigation induced a low oxygen concentration in soil microsites (Ju and Christie 2011; Zhang et al. 2017). However, in most cases the abundance of denitrification genes is probably not related to the  $\text{N}_2\text{O}$  flux (Cuhel et al. 2010; Liu et al. 2013). In the present study, the abundances of all detected nitrification and denitrification genes were not correlated with  $\text{N}_2\text{O}$  fluxes, indicating that the abundance of these genes might not be the limiting factor. Recent studies have shown that DNA-level analyses alone may not suffice to link soil microbial abundance with biogeochemical processes (Watanabe et al. 2009). Transcriptional analysis (mRNA) of functional genes might be applied to detect metabolically active organisms responsible for  $\text{N}_2\text{O}$  production in future studies. On the contrary, we show that the community compositions of nitrifiers and denitrifiers were more significantly related to  $\text{N}_2\text{O}$  fluxes than gene abundance, except for the AOA *amoA* gene. Why was the community composition of AOA not related to  $\text{N}_2\text{O}$  fluxes? It has been reported that AOAs are ubiquitous in soils but do not respond to  $\text{NH}_4^+$  oxidation or  $\text{N}_2\text{O}$  production in intensively managed agricultural soils (Di et al. 2009; Jia and Conrad 2009). We found that *Nitrosospira*-like AOBs were dominant over AOAs in  $\text{N}_2\text{O}$  production. Previous studies showed that  $\text{N}_2\text{O}$  emissions could be reduced by up to 77% using nitrification inhibitors of 3,4-dimethylpyrazol phosphate (DMPP) in the field, or by up to 80–99% using DMPP or dicyandiamide (DCD) in laboratory experiments (Ju and Christie 2011; Ma et al. 2015). Our results provide molecular microbial evidence that the community composition of nitrification and denitrification genes was a key factor for generating  $\text{N}_2\text{O}$  in the studied soils. These results also explain why nitrification or denitrification inhibitors were effective for reducing  $\text{N}_2\text{O}$  emissions in these soils. However, based on T-RFLP analysis, we still do not know the exact species diversity or the identities of

individual species that affect N<sub>2</sub>O production. High-throughput sequencing or cultivation-based approach might be used to identify the specific species and understand their roles in N<sub>2</sub>O production.

Compared to tree species, seasonal change had a weak effect on soil-microbe related N<sub>2</sub>O fluxes. For all detected nitrification and denitrification functional genes, the abundance of *narG*, *nirK*, and *nosZ* genes differed between seasons in soils of the oil-seed camellia forest, but not the slash pine forest. Community composition of the detected nitrifiers and denitrifiers varied slightly among sampling days. Although we selected seasonal sampling dates based on temperature, the variation in microbial community composition between soils was due to soil moisture, not to soil temperature. Soil moisture can enhance soil mineralization and denitrification rates: soil NH<sub>4</sub><sup>+</sup>-N and DOC contents increase through mineralization, while soil NO<sub>3</sub><sup>-</sup>-N concentrations decrease through denitrification (Friedl et al. 2016; Fujita et al. 2013;). Therefore, alteration of soil NO<sub>3</sub><sup>-</sup>-N, NH<sub>4</sub><sup>+</sup>-N, and DOC contents affects the microbial (gene) community composition of nitrifiers and denitrifiers, as we have shown. Szukics et al. (2010) found that both moisture and temperature affected nitrogen turnover rates by reversing the ratio of archaeal *amoA* to bacterial *amoA*. Phillips et al. (2014) found denitrification was more sensitive to moisture than to temperature, as moisture regulated the relative abundance of denitrifiers. A change in soil moisture implied that pH and oxygen concentrations in the soil might also be changed (DeAngelis et al. 2010). Therefore, soil conditions favouring denitrification would also rapidly alter the population of denitrifiers. Liu et al. (2012) reported that both abundance and composition of *narG*- and *nosZ*-containing denitrifiers responded to soil drying within only 1 day of surface water removal.

We found that in situ N<sub>2</sub>O flux was higher in soils of the oil-seed camellia forest in all seasons, than in slash pine forest soils. We found that the relative abundance of some T-RFs of all detected genes in soils of the oil-seed camellia forest significantly exceeded those in the soils of the slash pine forest, irrespective of season. Further, there were significant differences in soil moisture, NH<sub>4</sub><sup>+</sup>-N, and DOC contents between the two forest types. Therefore, we confirmed that tree species affected the soil environment in the study site, which affected the functional microbes, and ultimately induced difference in N<sub>2</sub>O flux in the soil.

## Conclusion

The difference in tree species, rather than between seasons, affected in situ N<sub>2</sub>O flux, soil physic-chemical factors, nitrifier and denitrifier abundance, and microbial (gene) community structure. In situ N<sub>2</sub>O flux was higher in soils of the oil-seed camellia forest than of the slash pine forest. The abundance and community composition of nitrifiers and denitrifiers was significantly affected by substrate availability, such as soil NO<sub>3</sub><sup>-</sup>-N and DOC contents. Community composition, but not abundance of nitrifier and denitrifier microorganisms, was related to N<sub>2</sub>O flux in acidic forest soils. Soil temperature, NO<sub>3</sub><sup>-</sup> and DOC concentrations, were significantly related to in situ N<sub>2</sub>O fluxes. We found that differences between tree species caused variation in the soil environment and in the community structure of functional microbes, and that ultimately these variations induced the different in situ N<sub>2</sub>O fluxes of the acidic forest soils. We suggest that in situ N<sub>2</sub>O emission from these soils is regulated by the balance between nitrifying and denitrifying microbial communities and their different biological processes, as well as the abiotic processes mediated by soil temperature changes over the year.

**Acknowledgments** This research was financially supported by the National Key Research and Development Program of China (2016YFD0200307), the National Natural Science Foundation of China (41771335, 41271280) and Hunan Provincial Natural Science Foundation of China (2016JJ3133). We would like to thank Editage [[www.editage.cn](http://www.editage.cn)] for English language editing.

## References

- Ambus P, Zechmeister-Boltenstern S, Butterbach-Bahl K (2006) Sources of nitrous oxide emitted from European forest soils. *Biogeosciences* 3:135–145
- Binkley D, Giardina C (1998) Why do tree species affect soils? The warp and woof of tree-soil interactions. *Biogeochemistry* 42:89–106
- Brown JR, Blankinship JC, Niboyet A, van Groenigen KJ, Dijkstra P, Le Roux X, Leadley PW, Hungate BA (2012) Effects of multiple global change treatments on soil N<sub>2</sub>O fluxes. *Biogeochemistry* 109:85–100
- Chen Z, Luo X, Hu RG, Wu MN, Wu J, Wei WX (2010) Impact of long-term fertilization on the composition of denitrifier communities based on nitrite reductase analyses in a paddy soil. *Microb Ecol* 60:850–861
- Cuhel J, Simek M, Laughlin RJ, Bru D, Chèneby D (2010) Insights into the effect of soil pH on N<sub>2</sub>O and N<sub>2</sub> emission

- and denitrifier community size and activity. *Appl Environ Microbiol* 76:1870–1878
- DeAngelis KM, Silver WL, Thompson AW, Firestone MK (2010) Microbial communities acclimate to recurring changes in soil redox potential status. *Environ Microbiol* 12:3137–3149
- Di HJ, Cameron KC, Shen JP, Winefield CS, O'Callaghan M (2009) Nitrification driven by bacteria and not archaea in nitrogen-rich grassland soils. *Nat Geosci* 2:621–624
- Friedl J, Scheer C, Rowlings DW, McIntosh HV, Strazzabosco A (2016) Denitrification losses from an intensively managed sub-tropical pasture – impact of soil moisture on the partitioning of N<sub>2</sub> and N<sub>2</sub>O emissions. *Soil Biol Biochem* 92:58–66
- Fujita Y, Bodegom PMV, Venterink HO, Han R, Witte JPM (2013) Towards a proper integration of hydrology in predicting soil nitrogen mineralization rates along natural moisture gradients. *Soil Biol Biochem* 58:302–312
- Guardia G, Abalos D, García-Marco S, Quemada M, Alonso-Ayuso M, Cárdenas LM, Dixon ER, Vallejo A (2016) Integrated soil fertility management drives the effect of cover crops on GHG emissions in an irrigated field. *Biogeosci Discuss* 13:5245–5257
- IPCC (2014) In: Pachauri RK, Meyer LA (Eds) *Climate change 2014: synthesis report. Contribution of working groups I, II and III to the fifth assessment report of the intergovernmental panel on climate change*, Geneva, Switzerland, p 151
- Jia Z, Conrad R (2009) Bacteria rather than archaea dominate microbial ammonia oxidation in an agricultural soil. *Environ Microbiol* 11:1658–1671
- Ju X, Christie P (2011) Calculation of theoretical nitrogen rate for simple nitrogen recommendations in intensive cropping systems: a case study on the North China plain. *Field Crop Res* 124:450–458
- Leininger S, Urlich T, Schloter M, Schwark L, Qi J, Nicol GW, Prosser JI, Schuster SC, Schleper C (2006) Archaea predominate among ammonia-oxidizing prokaryotes in soils. *Nature* 442:806–809
- Levy-Booth DJ, Prescott CE, Grayston SJ (2014) Microbial functional genes involved in nitrogen fixation, nitrification and denitrification in forest ecosystems. *Soil Biol Biochem* 75: 11–25
- Liu LL, Greaver TL (2009) A review of nitrogen enrichment effects on three biogenic GHGs: the CO<sub>2</sub> sink may be largely offset by stimulated N<sub>2</sub>O and CH<sub>4</sub> emission. *Ecol Lett* 12: 1103–1117
- Liu JB, Hou HJ, Sheng R, Chen Z, Zhu YJ, Qin HL, Wei WX (2012) Denitrifying communities differentially respond to flooding drying cycles in paddy soils. *Appl Soil Ecol* 62: 155–162
- Liu X, Chen CR, Wang WJ, Hughes JM, Lewis T, Hou EQ, Shen JP (2013) Soil environmental factors rather than denitrification gene abundance control N<sub>2</sub>O fluxes in a wet sclerophyll forest with different burning frequency. *Soil Biol Biochem* 57:292–300
- Luo HD, Zhu CF, Zhang L, Hu DN, Tu SP, Guo XM, Niu DK (2016) Effects of phosphorus fertilization levels on vegetative growth in *Camellia oleifera*. *Non-wood Forest Res* 34: 52–56
- Ma L, Shan J, Yan XY (2015) Nitrite behavior accounts for the nitrous oxide peaks following fertilization in a fluvo-aquic soil. *Biol Fertil Soils* 51:563–572
- Menyailo OV, Hungate BA (2015) Tree species and moisture effects on soil sources of N<sub>2</sub>O: quantifying contributions from nitrification and denitrification with 18-O isotopes. *J Geophys Res Biogeosci* 111(G2)
- Menyailo OV, Hungate BA, Zech W (2002) The effect of single tree species on soil microbial activities related to C and N cycling in the siberian artificial afforestation experiment: tree species and soil microbial activities. *Plant Soil* 242:183–196
- Morishita T, Aizawa S, Yoshinaga S, Kaneko S (2011) Seasonal change in N<sub>2</sub>O flux from forest soils in a forest catchment in Japan. *J For Res* 16:386–393
- Oura N, Shindo J, Fumoto T, Toda H, Kawashima H (2001) Effects of nitrogen deposition on nitrous oxide emissions from the forest floor. *Water Air Soil Pollut* 130:673–678
- Phillips RL, McMillan AMS, Palmada T, Dando J, Giltrap D (2014) Temperature effects on N<sub>2</sub>O and N<sub>2</sub> denitrification end-products for a New Zealand pasture soil. *N Z J Agric Res* 58:89–95
- Qin HL, Yuan HZ, Zhang H, Zhu YJ, Yin CM, Tan ZJ, Wu JS, Wei WX (2013) Ammonia-oxidizing archaea are more important than ammonia-oxidizing bacteria in nitrification and NO<sub>3</sub><sup>-</sup>-N loss in acidic soil of sloped land. *Biol Fertil Soils* 49:767–776
- Ravishankara AR, Daniel JS, Portmann RW (2009) Nitrous oxide (N<sub>2</sub>O): the dominant ozone-depleting substance emitted in the 21st century. *Science* 326:123–125
- Schindlbacher A, Zechmeister-Boltenstern S, Butterbach-Bahl SK (2004) Effects of soil moisture and temperature on NO, NO<sub>2</sub>, and N<sub>2</sub>O emissions from European forest soils. *J Geophys Res Atmos* 109:302
- Schleper C (2010) Ammonia oxidation: different niches for bacteria and archaea? *ISME J* 4:1092–1094
- Szukics U, Abell GCJ, Hodl V, Mitter B, Sessitsch A, Hackl E, Zechmeister-Boltenstern S (2010) Nitrifiers and denitrifiers respond rapidly to changed moisture and increasing temperature in a pristine forest soil. *FEMS Microbiol Ecol* 72:395–406
- Wang L, Sheng R, Yang HC, Wang Q, Zhang WZ, Hou HJ, Wu JS, Wei WX (2017) Stimulatory effect of exogenous nitrate on soil denitrifiers and denitrifying activities in submerged paddy soil. *Geoderma* 286:64–72
- Watanabe T, Kimura M, Asakawa S (2009) Distinct members of a stable methanogenic archaeal community transcribe *mcrA* genes under flooded and drained conditions in Japanese paddy field soil. *Soil Biol Biochem* 41:276–285
- Wolf I, Brumme R (2002) Contribution of nitrification and denitrification sources for seasonal N<sub>2</sub>O emissions in an acid German forest soil. *Soil Biol Biochem* 34:741–744
- Xu Y, Yang J, Chen W, Hou HJ, Li WH, Chen CL, Qin HL (2018) Bacterial community structure and abundance of plantation ecosystem soil response to seasonal change in red soil hilly region of southern. *J Sou Agr* 49:1289–1296
- Yang CD, Wu J, Tan HL, Liu YX, Xiong LM, Zhou LQ, Xie RL, Huang GQ, Zhao QG (2014) Comparison on soil microbial activities and bacterial diversity between *Betula alnoides* and *Pinus massoniana* plantations in red soil region. *Ecol Environ Sci* 23:415–422
- Yang O, Norton JM, Stark JM, Reeve JR, Habteselassie MY (2016) Ammonia-oxidizing bacteria are more responsive than archaea to nitrogen source in an agricultural soil. *Soil Biol Biochem* 96:4–15

- Zhang GB, Ji Y, Ma J, Xu H, Cai ZC, Yagic K (2012) Intermittent irrigation changes production, oxidation, and emission of CH<sub>4</sub> in paddy fields determined with stable carbon isotope technique. *Soil Biol Biochem* 52:108–116
- Zhang Y, Ji GD, Wang RJ (2016) Functional gene groups controlling nitrogen transformation rates in a groundwater-restoring denitrification biofilter under hydraulic retention time constraints. *Ecol Eng* 87:45–52
- Zhang ZX, Zhang WZ, Yang HC, Sheng R, Wei WX, Qin HL (2017) Elevated N<sub>2</sub>O emission by the rice roots: based on the abundances of *narG* and bacterial *amoA* genes. *Environ Sci Pollut Res* 24:2116–2125
- Zheng H, Ouyang ZY, Zhao TQ, Wang XK, Miao H, Peng YB (2006) Effect of different forest restoration approaches on soil biological properties. *Chin J App Environ Biol* 12:36–43
- Zheng XH, Mei BL, Wang YH, Xie BH, Wang YS, Dong HB, Xu H, Chen GX, Cai ZC, Yue J (2008) Quantification of N<sub>2</sub>O fluxes from soil–plant systems may be biased by the applied gas chromatograph methodology. *Plant Soil* 311:211–234